1 2	Coral Reef Ocean Acidification Monitoring: Development of a US Monitoring Strategy		
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13 14 15 16 17 18 19 20 21 22 23 24 25 26 27 28	 Introduction (Steering) Introduction (Steering) SOS CROAMP Workshop (Gledhill) Workshop Charge Workshop logistics, composition, format CROAMP Framework (Gledhill) Thematic Area → Strategic → Tactical → Observing Requirements Direct Chemistry Changes in Response to OA. (Gledhill/Andersson) Changes in biodiversity within coral reef ecosystems (Braindard) Community-scale impacts and feedbacks (Yates) Dissolution and bioerosion (Andersson) Organism response (Miller) Existing Monitoring Capacity (Sutton) 9.1. Report out on pre-workshop survey and capacity talks 10. Monitoring Expectations (Steering) 		
29	ABSTRACT		

The Coral Reef Ocean Acidification Monitoring Portfolio (CROAMP) Workshop was hosted by 30 the NOAA Ocean Acidification Program and the National Coral Reef Institute at the Nova 31 Southeastern University Oceanographic Center in August, 2012. Researchers and project 32 managers from around the world engaged in coral reef ocean acidification monitoring sought to 33 define a suite of metrics to optimally include as part of long-term coral reef monitoring. These 34 metrics comprise a listing of high-level strategic guiding questions towards which an observing 35 network should ideally be configured. These metrics are proposed as best able to discern 36 specific attribution of changes in coral reef ecosystems in response to ocean acidification. The 37 metrics necessarily constrain the observing requirements comprised of a suite of biogeochemical, 38 ecological, hydrologic, and meteorological measurements. We also clarify expectations for what 39 sustained ocean acidification monitoring can and cannot offer. Achieving the full CROAMP 40 requirements identified here will demand extensive leverage of existing national and 41 international initiatives particularly in an environment of increasing fiscal constraint. 42

43 I. INTRODUCTION

Ocean acidification poses many challenges to the long-term sustainability of coral reefs primarily 44 (although not exclusively) due to its capacity to reduce the rates of both biocalcification and 45 inorganic precipitation of carbonate minerals. Both individual organism and net community 46 calcification rates have generally been experimentally observed to decrease under expected 47 future OA conditions [e.g., Gattuso et al., 1998; Marubini et al., 2001, 2002; Marshall and Clode, 48 2002; Ohde and Hossain, 2004; Borowitzka, 1981; Gao et al., 1993; Langdon et al., 2000, 2003; 49 Langdon and Atkinson, 2005; Leclercq et al., 2000, 2002; Anthony et al., 2008]. Even under 50 51 contemporary conditions, field studies have revealed periods of net community dissolution [cite] with the expectation that these periods could increase under continued ocean acidification. Any 52

decline in net community calcification by coral reef organisms would likely compromise the
long-term persistence of many coral reef ecosystems because even the growth of healthy,
undisturbed coral reefs are known to only slightly outpace rates of loss due to physical and
biological erosion [see Glynn, 1997 for review].

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As US coral reef monitoring initiatives increasingly consider incorporation of ocean acidification 58 aspects into their designs, it's important that these investments be strategically directed towards 59 best informing our understanding of ecological consequences and feedbacks. Frequently, 60 monitoring studies offer only limited capacity to discern specific agents of change. Indeed, a 61 central challenge of field research is interpreting results amidst the sheer complexity of the 62 63 natural environment where change reflects the net result of multiple factors. Still, there remains a desire to engineer diagnostic monitoring most capable of testing response predictions of coral 64 reef ecosystems to ocean acidification as one means by which to identify potential refugia or 65 "hot spots" and while no metric has been identified which is exclusively sensitive to OA alone, 66 there is consensus that some are better than others. This report details those metrics and outlines 67 the measurement requirements necessary to derive them. 68

69

70 II. CROAMP WORKSHOP

A working group of nearly fifty cross-disciplinary multinational federal and academic scientists
was hosted by the NOAA Ocean Acidification Program and the National Coral Reef Institute at
the Nova Southeastern University Oceanographic Center, Center of Excellence for Coral Reef
Ecosystem Science, Dania Beach, FL on August 28-29, 2012. The workshop comprised 11

75 panel discussions detailing the current state-of-science and offering proposed monitoring metrics centered around five thematic areas: direct chemical changes, organism response, biodiversity 76 impacts, dissolution/bioerosion, and community-scale feedbacks. In addition, seven capacity 77 presentations were offered reviewing current and developing monitoring research efforts around 78 the world across multiple agency and academic efforts. A series of breakout sessions worked to 79 develop a portfolio of proposed metrics to be recommended for coral reef ecosystem monitoring 80 which would offer valuable insights into how coral reef ecosystem may be changing in response 81 to ocean acidification. 82



Figure1. Ocean acidification is anticipated to impact coral reefs in several ways. CROAMP
discussed monitoring for changes across five overarching thematic areas: A. detecting
direct chemical changes with the coral reef waters attributable to OA, B. detecting specific
organism responses to OA, C. tracking changes in biodiversity most susceptible to OA, D.
quantifying changes in dissolution and bioerosion rates, E. discerning changes and
feedbacks to whole-community metabolism.

The changing status of these metrics over time should aid in assigning specific attribution to OA, though it was recognized that there will likely be multiple factors driving observed changes we must be vigilant of co-variants. These workshop outcomes are intended to inform national and international long-term OA monitoring efforts within coral reef ecosystems.

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96 III. The CROAMP Framework

97 The CROAMP Framework was developed around the guiding principle that the observing
98 requirements should be derived from high-level overarching questions addressing how key
99 aspects of coral reef ecosystems might be anticipated to change over-time based upon the current
100 state-of-the-science. Within each thematic area, the following was developed:

Strategic Monitoring Goal (SMG) – These represent the high-level overarching goals towards
 which an observing portfolio is intended to answer. It indicates "*what*" we seek to know about
 the system with regards to its ultimate response to ocean acidification. Each strategic monitoring
 goal includes units of measure or is explicitly reported as unitless as applicable.

105 *Tactical Monitoring Metric* (TMM) – A TMM represents one or more specific tasks necessary to 106 achieve each of the strategic monitoring goals. These detail "*how*" we will acquire the requisite 107 information including the methods, associated measured parameters, and optimal frequency of 108 measure. As is the case with the Strategic Monitoring Goal, each Tactical Monitoring Metric 109 also includes units of measure. While various methods may be available to meet the tactical 109 requirements, the intent of this document is not to provide an exhaustive detail of all methods that could be employed, but simply to offer valuable examples. Other important considerations
in developing each of the TMM's included the feasibility of adoption into proposed monitoring
initiatives and what site criteria would be most amendable towards deriving the proposed TMM.
Finally, a critical aspect which was discussed as part of each TMM is the number and importance
of co-variants which represent factors other than OA which will convolute assigning specific
attribution. In nearly all cases, multiple co-variants were identified and are reflected in the
amount of secondary measurements which would need to be obtained simultaneously.



118

119 Figure 2. A guiding principle of CROAMP is that observing requirements are determined

120 by first carefully formulating tactical metrics that specifically inform strategic

121 requirements within each thematic area. Therefore, the primary emphasis of CROAMP

122 was upon developing the strategic requirements in a deliberate "top-down" fashion.



125 By far the most certain consequence of ocean acidification are the direct chemistry changes in response to rising atmospheric CO₂ concentration which are reflected in surface waters as 126 reduction in pH, an enrichment in total dissolved inorganic carbon, and a decrease in the 127 availability of carbonate ion. Such changes are already apparent in the extended time-series 128 observations at the Bermuda Atlantic Time-Series (BATS) and Hawaii Ocean Time-Series. 129 While changes in the hydrologic cycle, ocean circulation, and primary productivity may augment 130 these rates regionally over decadal time-scales, generally the seasonally detrended changes 131 observed in the surface waters (upper 50m) at these stations closely approximated near 132 equilibrium with rising atmospheric CO_2 . However, there is reason to suspect that such rates 133 may not precisely equate to the rates observed within many coral reef environments due to local 134 feedback and buffering processes. The relative influence local processes will have on the 135 overlay ocean acidification rate will be dependent upon several factors including water mass 136 residence time, benthic community composition, and net community metabolism. These 137 feedbacks will be considered more closely in the next section. 138

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A first-order question any ocean acidification monitoring observatory needs to be able to answer 140 is if the rates of ocean acidification are consistent with the regional oceanic average, or are they 141 decoupled. If they are decoupled then answering "why" will likely be addressed in the 142 subsequent thematic areas. To evaluate "if" they are decoupled demands that one monitoring not 143 only carbonate chemistry dynamics within the reef environment of interest, but also have careful 144 constraint on the rate of changes of neighboring oceanic waters where the relative influence of 145 146 coastal and benthic processes in negligible. This allows one to normalize the observed rates of change within the reef environment to the broader regional rates of change. 147

What is the annual rate of change in seawater carbonate chemistry within reef environments relative to comparitive rates of changes in neighboring open ocean surface waters? (Ω yr⁻¹, μ atm pCO_{2,sw} yr⁻¹, μ mol DIC kg⁻¹ yr⁻¹, μ mol TA kg⁻¹ yr⁻¹, pH_{Total} yr⁻¹)

Annual mean carbonate chemistry (CO2sys) within reef environment. (µmol kg⁻¹)

Primary Measured Parameters: Temp (°C), Salinity (unitless), TCO₂ (µmol kg⁻¹), TAlk (µmol kg⁻¹), pCO_{2,sea} (µatm)
Secondary Measured Parameters: pH (Total), pCO_{2,air} (µatm), Total P (µmol kg⁻¹), Total Si (µmol kg⁻¹)

Annual mean carbonate chemistry (CO2sys) within neighboring open ocean surface waters. (µmol kg⁻¹)

Primary Measured Parameters: Temp (°C), Salinity (unitless), TCO₂ (µmol kg⁻¹), TAlk (µmol kg⁻¹), pCO_{2,sea} (µatm)
Secondary Measured Parameters: Temp (°C), Salinity (unitless), TCO₂ (µmol kg⁻¹), TAlk (µmol kg⁻¹), pCO_{2,sea} (µatm)

Secondary Measured Parameters: PH (Total), pCO_{2,air} (µatm), Total P, Total Si

150

151 Figure 3. The CROAMP requirements for long-term monitoring of the direct changes in

152 *chemistry* in response to OA.

153 Estimates of the rate of change in aragonite saturation state (Ω_{arg}) within tropical surface waters

is on the order of 0.1 Ω_{arg} per decade. While such rates are readily discernible from multi-

decadal time-series stations within oceanic waters using sampling frequencies as coarse as

156 monthly, the variability within reef environments can be an order of magnitude greater

157 demanding much high sampling frequencies.TO ADD QUANTITATIVE ANALYSIS OF

158 TIME SERIES DATA TO EVALUATE MINIMUM FREQUENCY.

159 For monitoring within the reef environments, subdaily (3-hourly) is probably preferred to avoid

aliasing the data but bi-weekly discrete sampling is probably the minimum sampling frequency.

161 If adequate preliminary data is available, an appropriate frequency should be determined using a

Fourier series. The oceanic reference station will depend on the open ocean ecosystem chosenfor comparison, but in many cases, seasonal to monthly measures may be adequate.

164

Site criteria... To skillfully detect changes in response to OA it will be important to select sites 165 where local biogeochemistry is reasonably constrained and coastal processes are a minimal 166 influence. However, it is recognized that at least a few sites should be chosen to monitor such 167 changes within the context of local anthropogenic influences for comparative purposes. At those 168 sites, additional effort may be required to characterize influencing coastal processes. This 169 tactical monitoring index is best suited to fixed time-series stations as it is unlikely that stratified 170 sampling could be conducted at sufficient temporal frequency to avoid aliasing the data for 171 172 regression analysis. However, in systems of relatively lower short-term variability (e.g. low residence time systems) it may be possible to derive this index from stratified sampling provided 173 it was maintained for several decades. 174 co-variants - Most of these measures will co-vary with temperature, salinity, and any processes 175

176 which might alter the CO2SYS.

177

178 *Methods:*

Provided a measure of any two of these four parameters, the other two can theoretically be
solved. To do so, the CO2 System needs to be solved using selected dissociation constants. The
carbonate equilibria calculations should use the Mehrbach et al. [1973] formulations of the K1
and K2 dissociation constants as refit by Dickson and Millero [1987]. The sensitivity with

183	which the saturation states and CO2-3 can be calculated is dependent on the parameter
184	measured; pH and pCO2 yield the greatest uncertainty, followed by the TA, DIC pair. While
185	triple constraint is advised within coastal systems, if only two parameters are to be measured, pH
186	or pCO2 with DIC yields the greatest precision, closely followed by pH or pCO2 and TA. All
187	measurements should follow procedures and protocols outlined in the Best Practices guides
188	(Dickson et al., 2007; Riebesell et al., 2010).

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191 2. Community Metabolism Response and Feedbacks to Ocean Acidification

Changes in community-scale processes including net community productivity (P), respiration
(R), calcification (G), and dissolution (D) can significantly impact the local surrounding
chemistry thereby overprinting the affects of OA in some cases. Understanding the interactions
between coral reef communities and the surrounding chemical environment is critical towards
improving our understanding with regards to how OA is unfolding within these systems and
what local processes might prove dominant.

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Ocean acidification may profoundly affect the basic ecological interactions which structure coral reef ecosystems, providing even greater challenges for local management strategies aimed at retaining or regaining coral dominance. A recent meta-analysis of available experimental data

- suggests that such 'phase shifts' might increase in response to OA. Specifically, as calcification
- is reduced, algae (non-calcifying) and sea grass may benefit [Hendriks et al., 2010].

What are the long-term trends in annual integrated net community metabolic performance? (e.g. mol C m⁻² y⁻¹, mol CaCO3 m⁻² y⁻¹, NEP:NEC (C/CaCO3), mm or cm yr⁻¹)



205

- 206 Figure 4. The CROAMP requirements for long-term monitoring *Community Metabolism*
- 207 *Response* to OA.

- An alternative approach to paired TA and DIC measures might be tracking shifts in the
- regression slope of pH vs O2. While simpler to implement than the paired measurements, this

approach remains largely speculative at this point and needs to be better documented. Good
quality autonomous instruments are increasingly available for short duration (e.g. weeks to a few
months) deployments. High quality pH measures are best obtained through discrete sample
collection according to best practices. Spectrophotometric determination can then be performed
post-collection. Discrete samples need to be measured within an hour of collection.

216

Site criteria - To skillfully detect community changes in response to OA it will be important to
select sites where local biogeochemistry is reasonably constrained and coastal processes are a
minimal influence. However, it is recognized that at least a few sites should be chosen to
monitor such changes within the context of local anthropogenic influences for comparative
purposes.

222

223 3. Dissolution and Bioerosion Response and Feedbacks to Ocean Acidification

Some species exhibit a natural tolerance or local environmental conditions can compensate for
global OA effects. It has now been well established that coral calcification rates are, in fact, a
function of multiple factors including light, temperature, carbonate chemistry and nutrient
supply.

Beyond the potential direct effects for coral calcification, OA poses additional concerns for coral
reef ecosystems. These include dissolution of coral reef sediments which often contain
appreciable amounts of more soluble carbonate minerals [Morse et al., 2006]. Sediment
dissolution may outpace carbonate production on many reefs by 2030 [Yates and Halley, 2006].

Furthermore, current-day areas where upwelling causes chemistries comparable to future OA 232 conditions (i.e. a tripling of atmospheric CO2) exhibit poorly cemented and highly bioeroded 233 coral reefs [Manzello et al., 2008]. These effects may compromise coral reef framework integrity 234 235 and resilience in the face of other acute threats such as coral bleaching, diseases, increases in storm intensity, and rising sea level [e.g., Silverman & Calderia, 2009]. Indeed, in CO2 enriched 236 waters of the Galapagos Islands, reef structures were completely eroded to rubble and sand in 237 less than 10 years following one acute warming disturbance [i.e., 1982-83 El Nino event: 238 Manzello, 2009]. 239

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What are the long-term trends in mechanical breakdown/dissolution rates of $CaCO_3$? (g $CaCO_3$ m⁻² yr⁻¹)

Annual mean alkalinity anomalies relative to neighboring open ocean surface waters ($\mu mol \ kg^{-1}$).
• Primary Measured Parameters: Temp (°C), Salinity (unitless), TAlk (μmol kg ⁻¹)
• Secondary Measured Parameters: Total P (μmol kg ⁻¹), Total Si (μmol kg ⁻¹
Bioerosion rates at specific sites. (g CaCO ₃ m ⁻² yr ⁻¹)
• Primary Measured Parameters: Bioeroder colonization density (number m ⁻²) and rate (number m ⁻² yr ⁻¹) of experimental substrate
• Secondary Measured Parameters: Epibenthic bioeroder density (individuals m ⁻²), Bioeroder size frequency distribution (number of size classes m ⁻²)
Metabolic CaCO ₃ dissolution rates. (g CaCO3 m ⁻² yr ⁻¹)
• Primary Measured Parameters: Temp (°C), Salinity (unitless), TAlk sediment profiles (µmol kg ⁻¹), TAlk sediment-seawater flux (µmol m ⁻² day ⁻¹), Mass loss (g CaCO ₃) of experimental substrate
• Secondary Measured Parameters:Light (moles photons m ⁻² s ⁻¹), nutrients (mmol kg sw ⁻¹), total organic carbon (TOC, mg m ⁻³)
Architectural complexity. (m m ⁻⁺)
Primary Measure Parameters: Rugosity (m m ⁻¹)
Secondary Measured Parameters:
Sediment mineralogy/composition (mol %)
Sediment mineralogy/composition (mor //)
 Primary Measured Parameter: carbonate mineral absolute (g CaCO3 g⁻¹ sediment) and relative abundance (calcite:aragonite:high-MgCO₃), mol fraction MgCO₃ (%)
• Secondary Measured Parameter: grain size distribution (# m ⁻²)

Figure 5. The CROAMP requirements for long-term monitoring of the *Dissolution and Bioerosion Response* in response to OA.

A measure of small-scale variations or amplitude in the height of a reef surface. Rugosity is a 244 measure of architectural complexity which is presumed to be an indicator of the amount of 245 habitat available for colonization by benthic organisms, and shelter and foraging area for fish. 246 Tracking rugosity over time can be a useful metric to examine structural changes in reefs as the 247 framework building corals die, rates of bioerosion overtake accretion, and the reefs eventually 248 flatten [18] although care must be taken to account for acute disturbance events (e.g. storm 249 effects and mass wasting) when assigning attribution to long-term trends. Structural complexity 250 is classically measured with a reef rugosity index, which is the ratio of a straight line transect to 251 the distance a flexible chain of equal length travels when draped over the reef substrate, however 252 recent advances in digital reef rugosity (DRR) have greatly reduced that labor required to obtain 253 a rugosity measure while making it a more objective and repeatable measure useful in statistical 254 analysis. Dustan P, Doherty O, Pardede S (2013) Digital Reef Rugosity Estimates Coral Reef 255 Habitat Complexity. PLoS ONE 8(2): e57386. doi:10.1371/journal.pone.0057386 256



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259 4. Organism Response to Ocean Acidification

260 Other expected impacts of OA include a potential lowering of the thermal thresholds for bleaching [Anthony et al., 2008], the impairment on early life stages of corals such as reduced 261 fertilization success, reduced larval settlement, and reduced growth and survival rates of newly 262 settled corals [Albright et al., 2008; Cohen and Holcomb, 2009; Albright et al., 2010, Morita et 263 al., 2010; Suwa et al., 2010]. These species-level impacts are particularly relevant given the 264 potential extension of Endangered Species Act protection for additional scleractinian coral reef 265 species as this Act provides for potent conservation actions for species at risk of extinction, and 266 impacts on population bottlenecks such as recruitment may pose more direct extinction risk then 267 268 calcification and growth impacts.

269

The affects of OA on some non-calcifying organisms may indeed be just as serious as the affectsto calcifiers but are largely unexplored.

272

In addition, predictions of coral reef ecosystem responses to OA are further complicated due to
the fact that reef organisms secrete species-specific types of calcium carbonate mineralogies (i.e.,
aragonite, calcite, and magnesium calcite) which exhibit a range of solubilities.

What are the long-term regional changes in growth rates of target taxa in response to OA?

Growth rates of select calcifiers.

Primary Measured Parameters: Coral skeletal linear extension rate (mm yr⁻¹), density (g cm⁻³), and mass (g)
Secondary Measured Parameters: Halimeda spp. (or other calcified algal taxa) growth rate (g yr⁻¹), Other taxa (e.g. forams, urchins) growth rate (g yr⁻¹)

Growth and productivity rates in select non-calcifiers. (g colony⁻¹ yr⁻¹)

Primary Measured Parameters: seagrass spp. plant mass (g), shoot density (# m⁻²), and photosynthetic rate (μg C μg chlorophyll σ⁻¹ hr⁻¹); fleshy algal spp. plant mass (g) and photosynthetic rate (μg C μg chlorophyll σ⁻¹ hr⁻¹)
 Secondary Measured Parameters: Temp (°C), Light (moles photons m⁻² s⁻¹), nutrients (mmol kg sw⁻¹)

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What are the long-term regional changes in 'recruitment' rates of target taxa in response to OA.

Crustose coralline algae accumulation rate (CCA, mm yr⁻¹).

Primary Measured Parameters: CCA accumulation rate on standardized substrates (mm yr-¹)
 Secondary Measured Parameters: Temp (°C), Light (moles photons m⁻² s⁻¹), nutrients (mmol kg sw⁻¹), current speed (cm sec⁻¹), grazing rate (g grazer⁻¹ d⁻¹)

Calcifying epiphytes accumulation rate. (μ g Chl *a* cm⁻² yr⁻¹)

Primary Measured Parameters: seagrass epiphyte density (μg Chl a cm⁻²)
Secondary Measured Parameters: Temp (°C), Light (moles photons m⁻² s⁻¹), nutrients (mmol kg sw⁻¹), current speed (cm sec⁻¹), grazing rate (g grazer⁻¹ d⁻¹)

What are the long-term regional changes in physiological 'condition' of individual organisms response to OA?

Plan carbon/nutrient status. Primary Measured Parameters: Seagrass tissues C:N (moles C moles N⁻¹), Algal tissues C:N (moles C moles N⁻¹) Secondary Measured Parameters: Temp (°C), Light (moles photons m⁻² s⁻¹), nutrients (mmol kg sw⁻¹), current speed (cm sec⁻¹) Organism mineral content. Primary Measured Parameters: calcifier structural minerology (polymorph) Secondary Measured Parameters: Temp (°C), Salinity (unitless), TCO₂ (µmol kg⁻¹), TAlk (µmol kg⁻¹), pCO_{2,sea} (µatm) Coral tissue thickness. Primary Measured Parameters: tissue thickness (mm) Secondary Measured Parameters: Temp (°C), Light (moles photons m⁻² s⁻¹), nutrients (mmol kg sw⁻¹), current speed (cm sec⁻¹)

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Figure 6. The CROAMP requirements for long-term monitoring of the Organism Response
in response to OA.

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283 5. Changes in Coral Reef Biodiversity to Ocean Acidification

While coral reefs are considered the most biologically diverse of all marine ecosystems this
diversity varies greatly across spatial and environmental gradients. How this variation influences
reef resilience and adaptation to climate change and OA is poorly understood. Much of the
biomass and biodiversity of reefs lies within the complex architecture of the reef matrix
[Ginsburg, 1983; Small et al., 1998; Sala and Knowlton, 2006; Knowlton et al., 2010]. The
potential effects of OA on this community of organisms, collectively known as the cryptobiota
[Macintyre et al., 1982], remains largely unknown.

- 292 Understanding the biodiversity and community structure of coral reef ecosystems is therefore
- 293 necessary in recognizing and predicting shifts in community structure in response to OA.

What are the long-term trends in biodiversity as appearent from changes in benthic composition, community structure, and ecological function.?



- **Figure 7. The CROAMP requirements for long-term monitoring of the** *Coral Reef*
- 299 *Biodiversity* response to OA.
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- 301 IV. Existing Ocean Acidification Monitoring Capacity
- 302 V. Existing Ocean Acidification Monitoring Capacity
- 303 VI. Summary
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- the ability and rate of calcium carbonate production by many marine calcifiers is anticipated to
- decrease while rates of bioerosion and dissolution could increase. Such affects would have
- important consequences for the role and function of coral reefs and compromise the ecoservices
- they afford us.
- 313 **References.**
- 314 Tables.

315 Summary of Measurements

Measurement (units)	Description	Methods	References	
Physicochemical				
Nutrients (mmol kg sw ⁻¹) –	These include all dissolved inorganic nutrients (ammonium, silicate, nitrate, nitrite, and phosphate).	Discrete water sample collection followed by post- collection analysis according to best-practices.		
Oxygen (µmol kg-1) –		Discrete water sample collection followed by post- collection analysis according to best-practices.		
pCO2,air (µatm) —				
pCO2,sw (µatm) –				
pH (Total) –				
Salinity (unitless) —	Salinity is a fundamental measurements for interpreting carbonate chemistry and must be obtained contemporaneously with any measurement of CO2SYS.			
TAlk (µmol kg-1) –				
TCO2 (µmol kg-1) –				
Temperature (oC) –	Temp. is a fundamental measurements for interpreting carbonate chemistry and must be obtained contemporaneously with any measurement of CO2SYS. Furthermore, it imparts a fundamental control on biology and thus needs to be a key measurement ecological monitoring.	Autonomous temperature sensors (e.g. CTD). Do not rely solely on manufacturers' stated accuracy; regular lab calibrations and atsea comparisons are recommended		
3) –				

Total P (µmol kg sw-1) –				
Total Si (µmol kg sw-1) –				
Ecological				
[Coral, algae, urchin and				
algae] absolute (OTU #) and				
relative abundance (%) –				
[Coral, algae, urchin and algael benthic cover (%) –				
<i>ICoral, algae, urchin and</i>				
algae/ diversity index				
(unitless) –				
[Coral, algae, urchin and				
algae] ID/richness (OTU,				
Organismal Taxonomic Unit) –				
[Coral, algae, urchin and				
algae] percent cover (%) –				
[Coral, algae, urchin and				
algae]density (# m ⁻²) –				
[Coral, algae, urchin and				
algaejevenness (unitiess) –				
moles N-1) –				
Bioeroder colonization density				
(number m^{-2}) –				
Bioeroder colonization rate				
$(number m^2 yr^2) -$				
Bioeroder size frequency				
distribution (number of size classes m^{-2})				
Biomass –				
Calcifier structural				
Curcipici siruciurui				

minerology (polymorph) –		
Coral skeletal density (g cm ⁻³) –		
Coral skeletal linear extension $rate (mm m^{-1})$		
rule (mm yr) –		
Corai skeletat mass $(g) =$		
Coral tissue thickness (mm) –		
Crustose coralline algae accumulation rate (CCA, mm yr ⁻¹) –		
Density of individuals (per m ²) -		
<i>Epibenthic bioeroder density</i> (individuals m^{-2}) –		
Fleshy algal spp. Photosynthetic rate ($\mu g C \mu g$ chlorophyll $a^{-1} hr^{-1}$) –		
Fleshy algal spp. plant mass (g) –		
Grazing rate (g grazer ⁻¹ d^{-1}) –		
Halimeda spp. (or other calcified algal taxa) growth rate (g yr ⁻¹) –		
Identification –		
Mass loss of experimental substrate (g CaCO3) –		
Other taxa (e.g. forams, urchins) growth rate (g yr ⁻¹) –		
Percent cover (%) –		
Seagrass epiphyte density (μg Chl a cm ⁻²) –		

	amplitude in the height of a reef surface. Rugosity is a measure of architectural complexity which is presumed to be an indicator of the amount of habitat available for colonization by benthic organisms, and	(DRR)	
Rugosity (STD depth, m) –	A measure of small-scale variations or	digital reef rugosity	Dustan et al., (2013)
(²) – Mole fraction MgCO ₃ (%) –			
Grain size distribution (# m ⁻			
Depth to sediment surface (m) –			
Current speed (m sec ⁻¹) –			
Community water depth (m)			
abundance (calcite:aragonite:high- MgCO3) –			
Carbonate mineral relative			
Carbonate mineral absolute abundance ($g CaCO_3 g^{-1}$			
	Hydrodynamic and Structural		
Substrate type –			
Seagrass tissues C:N (moles C moles N^{1}) –			
Seagrass spp. shoot density (# m^{-2}) –			
Seagrass spp. plant mass (g) -	-		
rate ($\mu g C \mu g$ chlorophyll a^{-1} hr^{-1}) –			
Sagarass snn nhotosynthetic			

Sediment thickness (cm) –			
TAlk sediment profiles			
$(\mu m ol k \sigma^{-1}) =$			
TAlk sediment-seawater flux			
$(umol m^{-2} dav^{-1}) -$			
		1	
Meteorological			
Light (moles photons $m^{-2} s^{-1}$)			
-			
Windspeed (m sec ⁻¹) –			

